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Understanding the formation of black soils

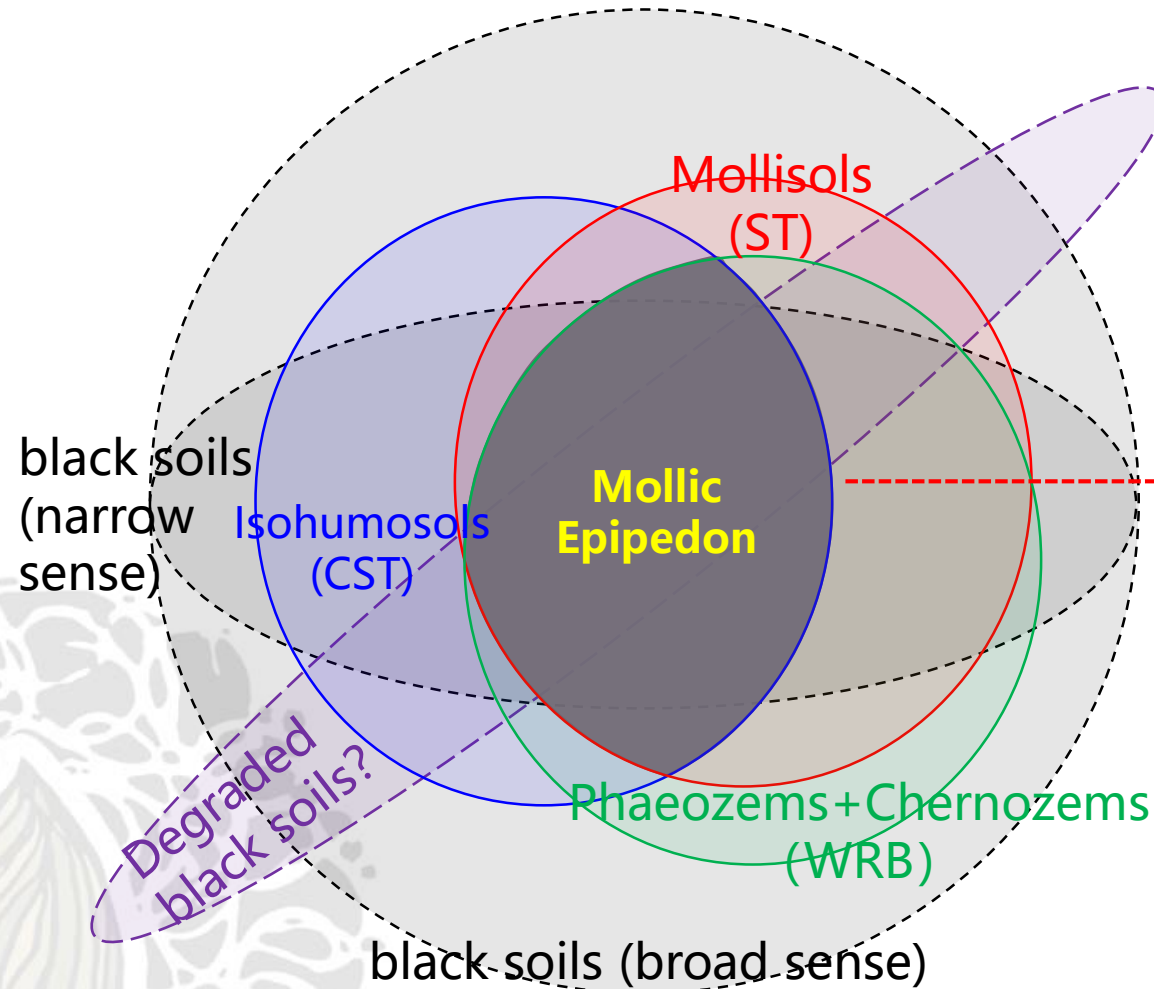
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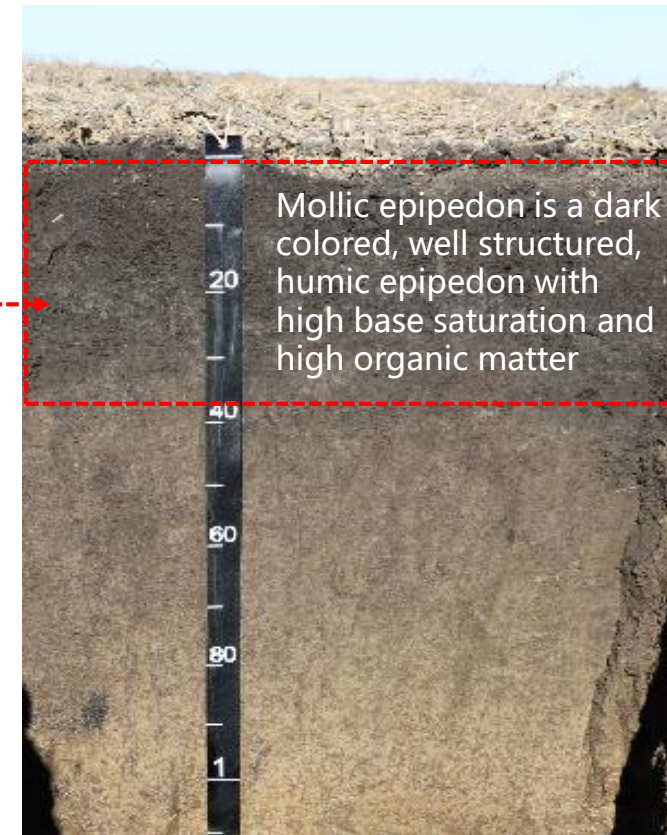
Webinar series
SUSTAINABLE
MANAGEMENT
OF BLACK SOILS



Blacks soils in different classification systems



Mollic epipedon is the core of black soils, and the basis for its high fertility

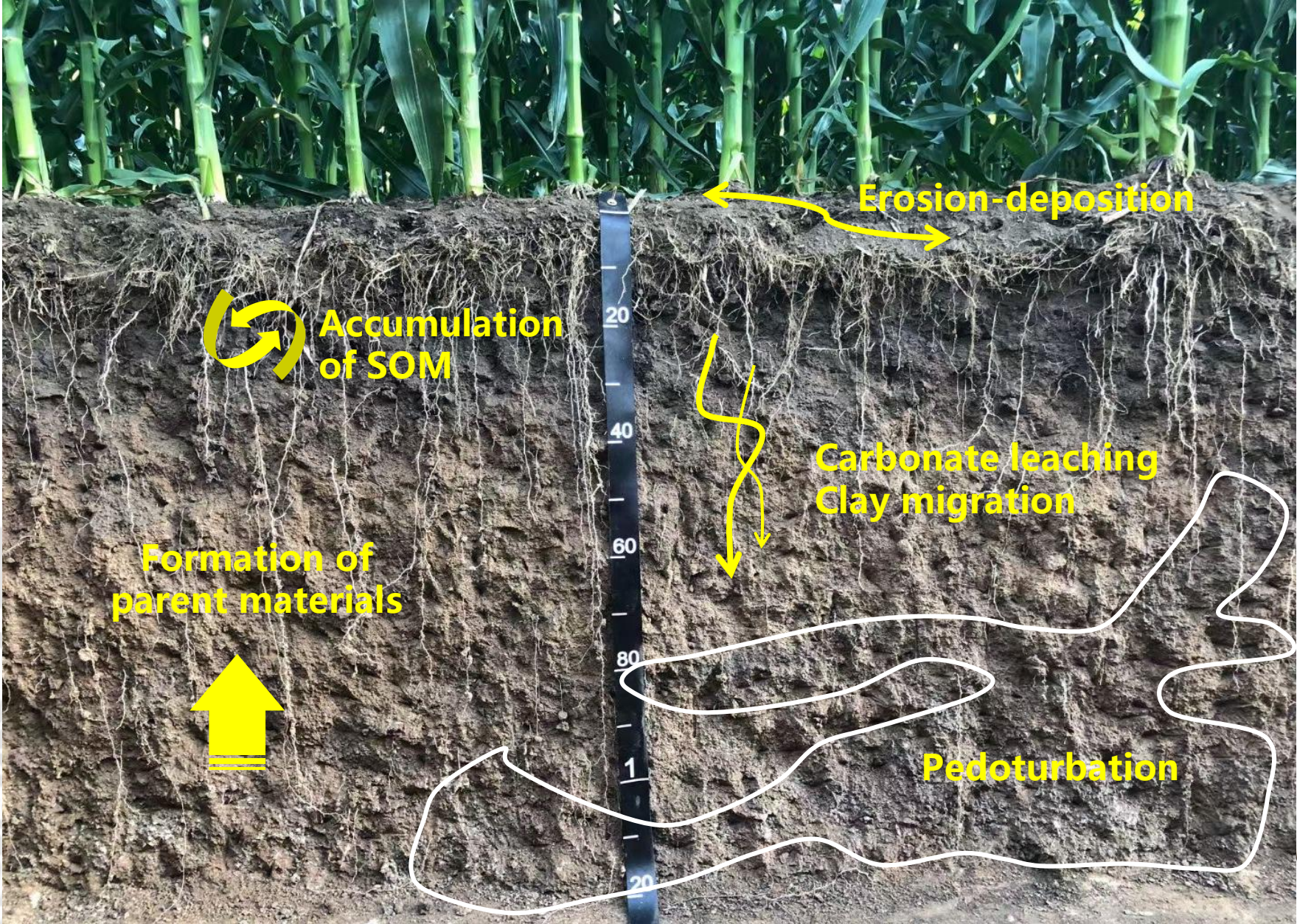


Blacks soils in different classification systems

Webinar series | SUSTAINABLE MANAGEMENT OF BLACK SOILS



Key processes in the formation of black soils



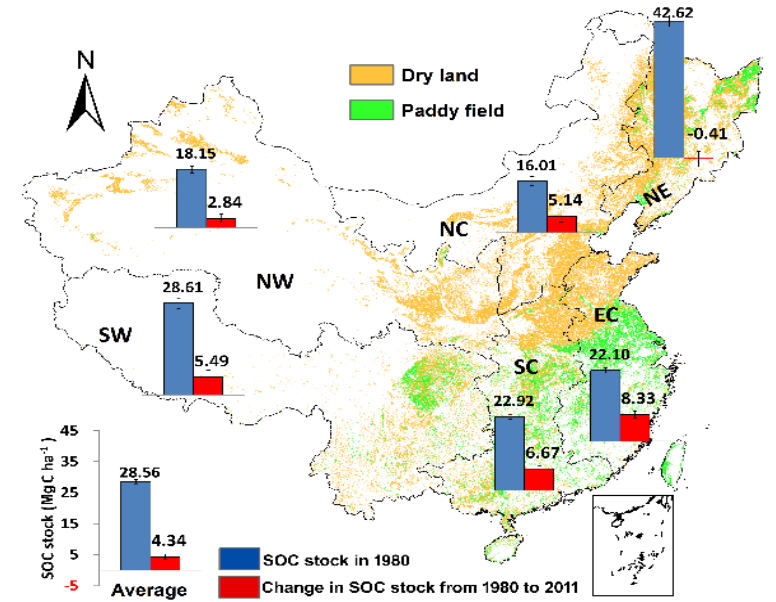
Black soils are under threats

Soil erosion



Two major threats of black soils

Loss of SOM



(Zhao et al., 2018, PNAS)

soil compaction



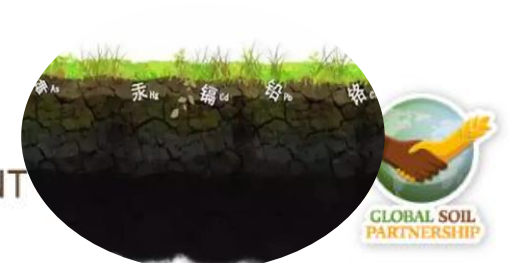
water logging



biodiversity loss



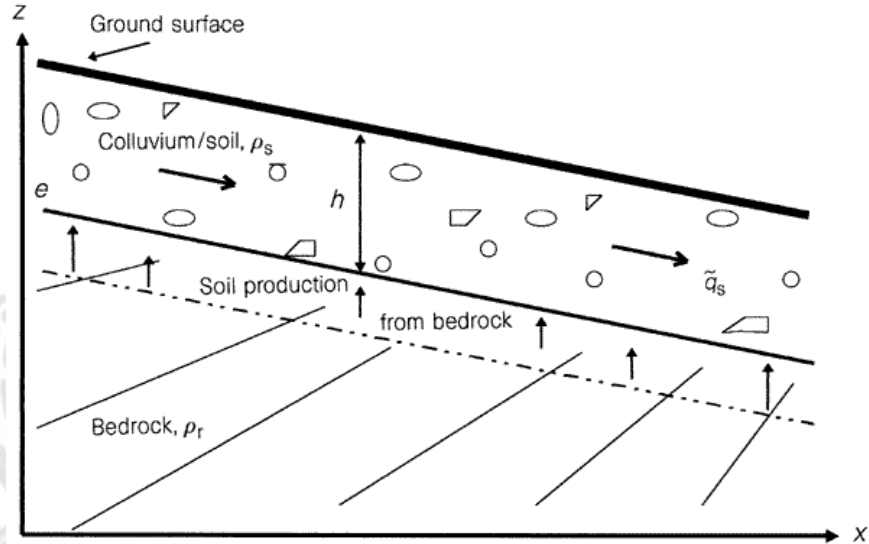
Soil pollution



Other threats

Sustainability of black soils

The sustainability of black soils depends on the balance between the rates of soil loss and the rates of soil formation.



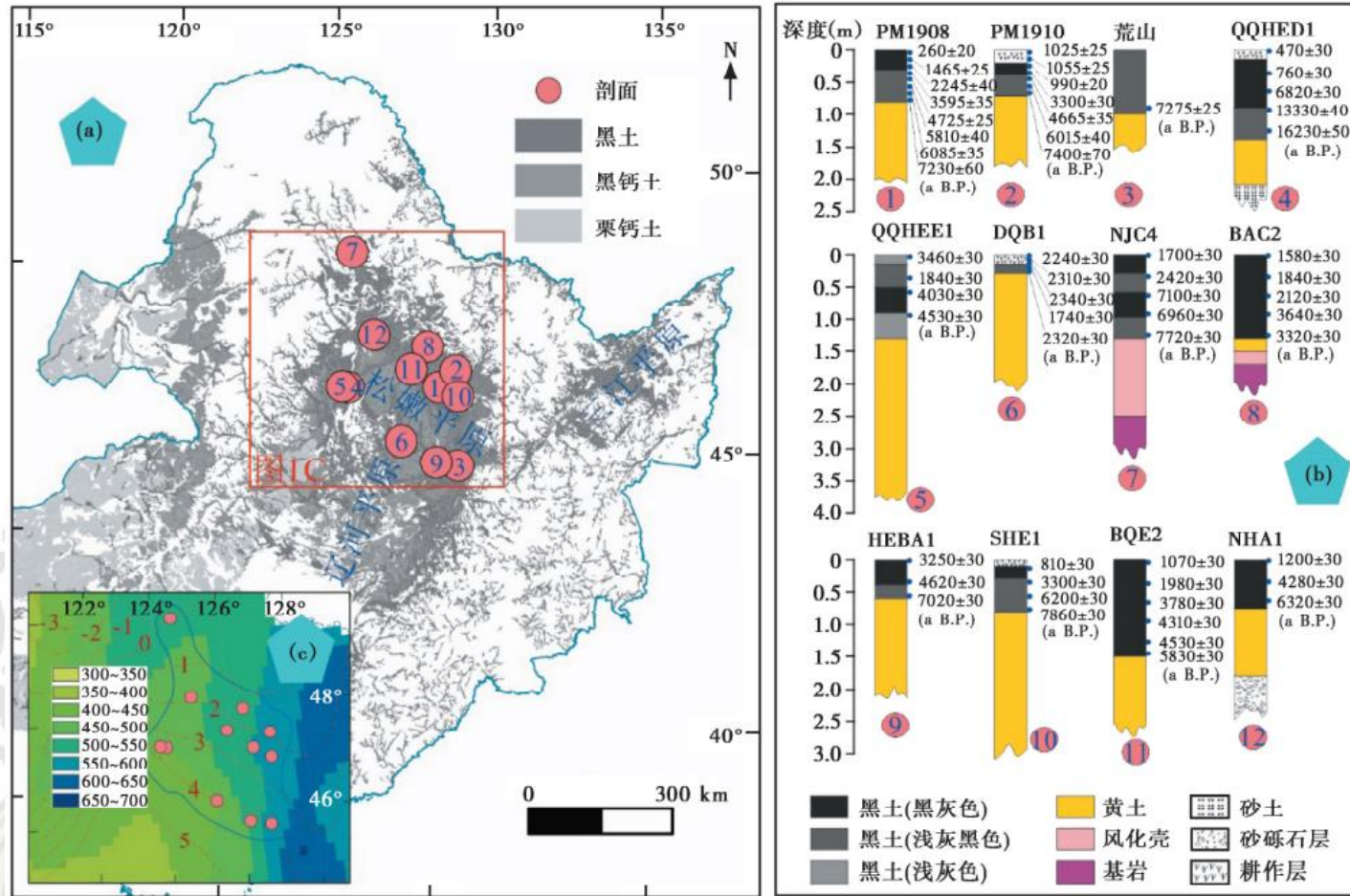
(Heimsath et al., 1997, Nature)

Soil thickness = soil formation (e.g., rock weathering, exogenous input) – soil loss (e.g., erosion, denudation)

- Black soils loss at average rates $0.2-0.3\text{cm yr}^{-1}$ (Zhang & Liu, 2020)
- How long does it take to form 1cm black soils?
 - 20-40 yrs (Lin et al., 1999)
 - ~100 yrs (Song et. al., 2020)
 - 120-400 yrs (Yan & Tang, 2005)
 - 300-400 yrs (Liu et al., 2003)

Knowledge on the rates of black soil formation is limited by the lack of precise soil chronology!

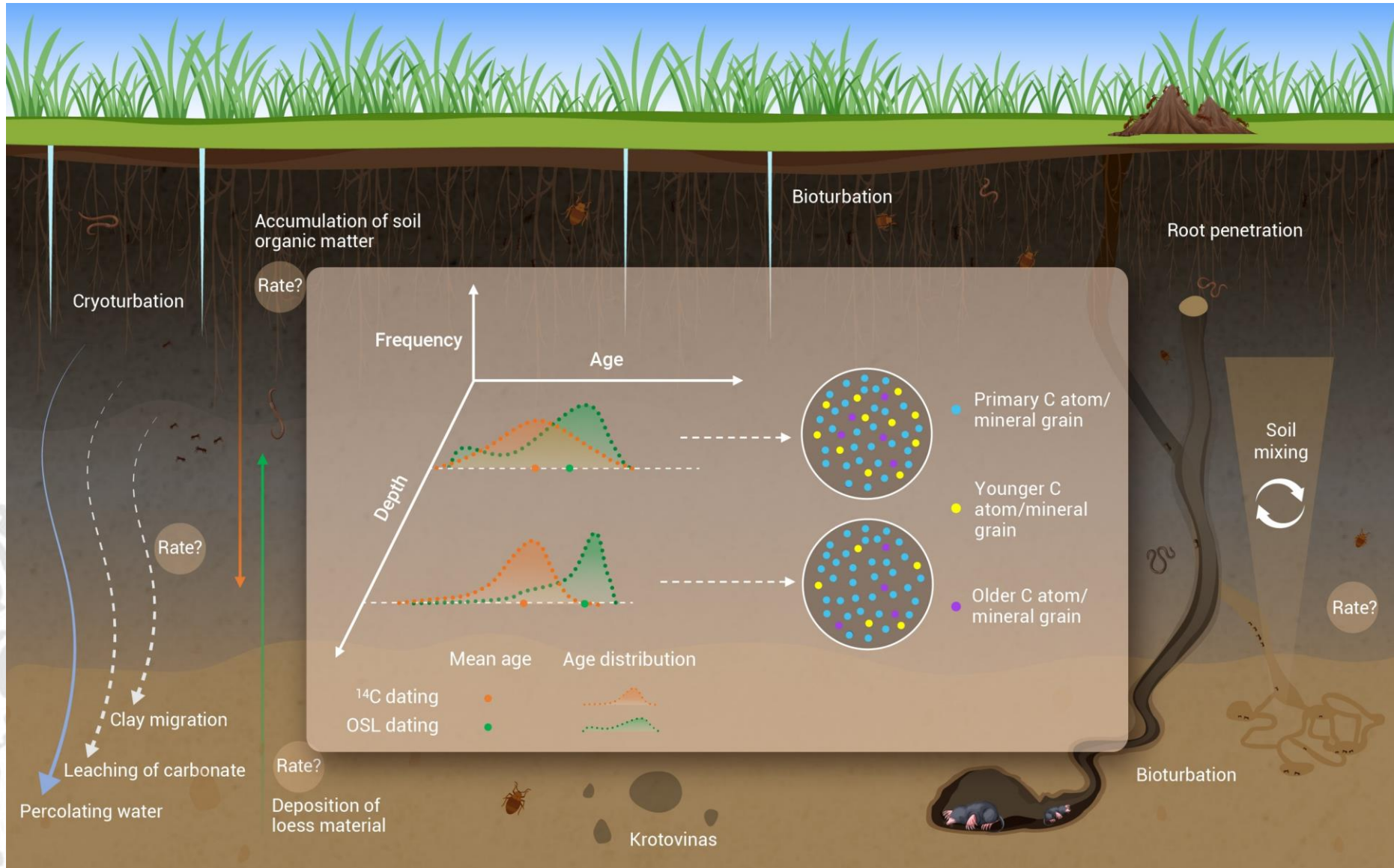
Previous knowledge on black soil ages



(Cui et al., 2021)

- Black soils in Northeast China formed during **early-mid Holocene**, based on ¹⁴C dating of SOM (Cui et al., 2021)
- Results from other regions of the world (e.g., Mid-USA, Europe, Argentina) also point to a **Holocene age** of black soils. Only in rare cases with an older age.

An open and complex system – dating soil is a challenge



Source of younger C:

- ① continuum nature
- ② rooting
- ③ leaching of DOC
- ④ pedoturbation

Source of older C:

- ① inherent
- ② erosion-deposition

Source of younger grain:

- ① pedoturbation
- ② clay migration

Source of older grain:

- ① erosion-deposition
- ② bioturbation
- ③ insufficient bleaching



MENT OF BLACK SOILS

Two protocols to determine black soil ages

Protocol 1

To find closed/ semi-closed systems,
such as charcoal, or buried soils

Protocol 2

To identify signals of primary age and pedoturbation,
fine-grained soil dating, such as single-grain OSL dating and ^{14}C
dating of SOM at molecular level

Upper age limit for black soils – charcoal



- A piece of woody charcoal beneath the Mollic layer aged **16.9 ka. a BP**, setting the upper age limit of black soils in northeast China.
- It could be forest landscape prior to the presence of black soils



		¹⁴ C age
	charcoal	16897±161 a BP
85cm	SOM	12742±19 a BP
	Humin	12992±100 a BP
	Humic acid	15870±150 a BP

Mollisols did not form at the time (16.9 ka BP) of the formation of the charcoal



Lower age limit for black soils – buried mollic layer



- ❑ Buried Mollic layer, fast erosion – deposition protect SOM from deposition
- ❑ ^{14}C age of the buried layer = C age prior to burial + time elapse since the burial



^{14}C age

	SOM	11873 ± 129 a BP
215cm	Humin	12287 ± a BP
	Humic acid	12548 ± a BP
	SOM	11835 ± a BP
275cm	Humin	12571 ± a BP
	Humic acid	12071 ± a BP

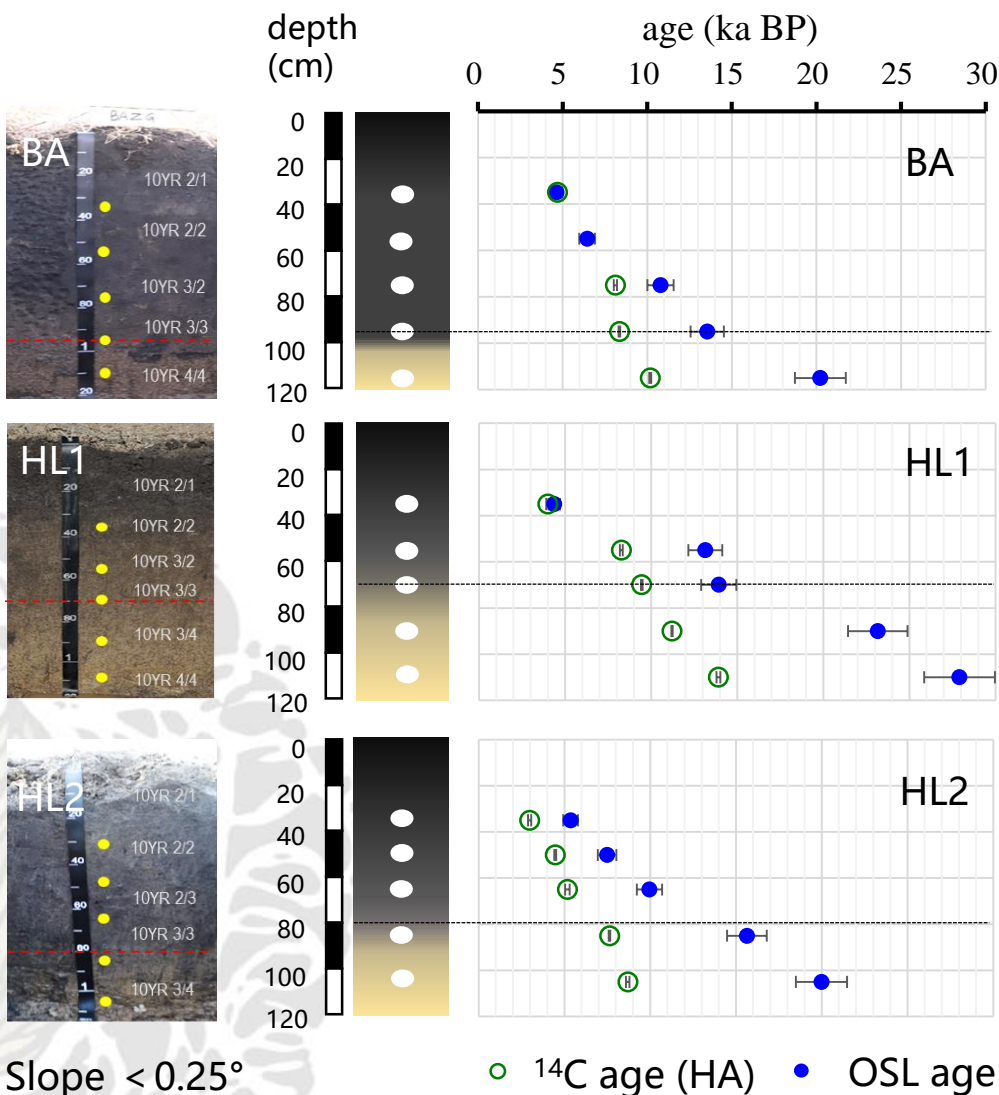
Well-developed black soils existed prior to 12.6 ka BP

Thus, black soils in northeast China started to form between 16.9-12.6 ka BP.



Ages discrepancies between ¹⁴C and OSL dating

¹⁴C and OSL dating of three soils from flat terrain



Slope < 0.25°

Lower boundary of mollic layers

¹⁴C ages ~ 7-9 ka BP

OSL ages ~ 13.5-14 ka BP

Bølling-Allerød warm period (14.7~12.7 ka BP)

Previous ¹⁴C dating of SOM may have underestimated the age of black soils.

➤ OSL ages are much older than ¹⁴C ages, especially in deeper soils.

➤ The OSL ages at the bottom of the mollic layer 13.5-14 ka BP, agree well with aforementioned results.

➤ Contamination by younger C from roots exudates, DOC leaching, pedoturbation...



Favorable environment for black soils formation

Vegetation expand in response to abrupt warming during the last deglaciation (Xu et al., 2023)

Temperature raised since 20 ka indicated by GDGT (Zhu et al., 2021)

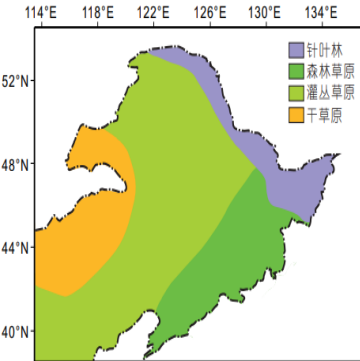
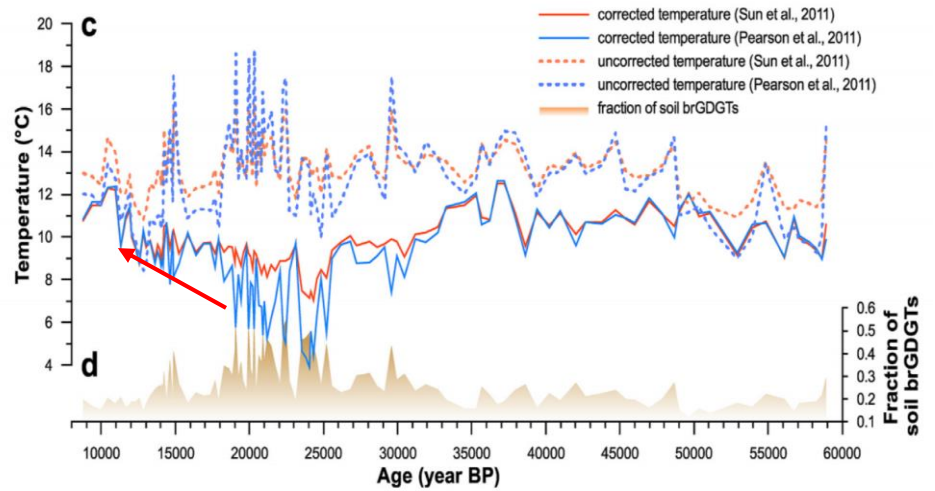
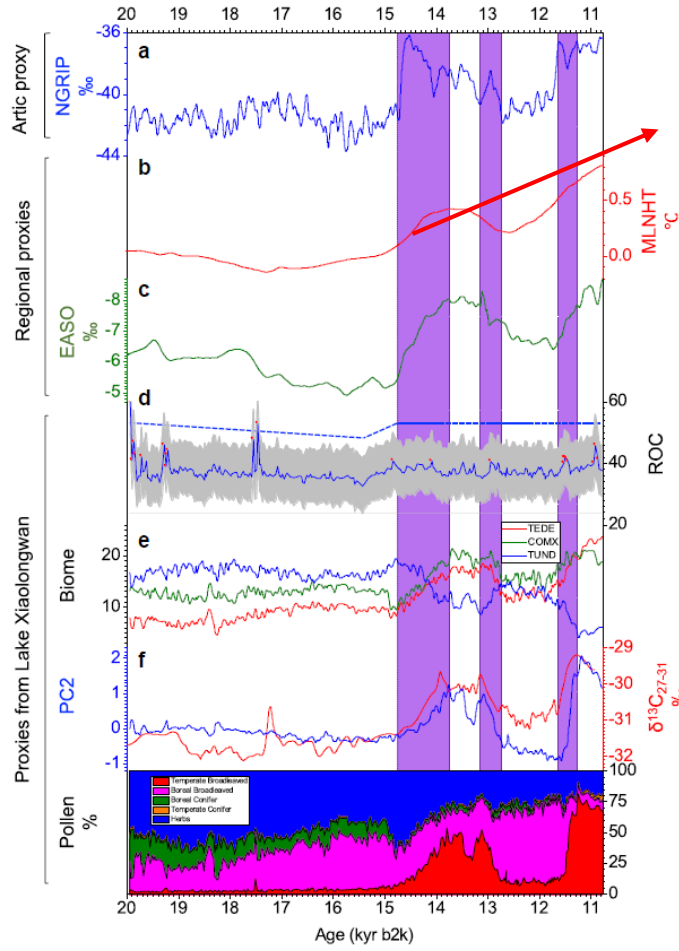


图 15 末次盛冰期东北地区植被格局

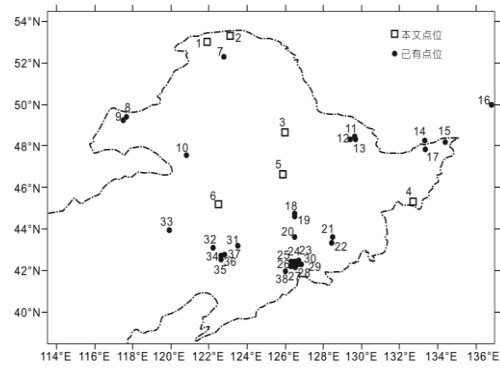


图 14 东北地区末次盛冰期以来花粉资料点分布

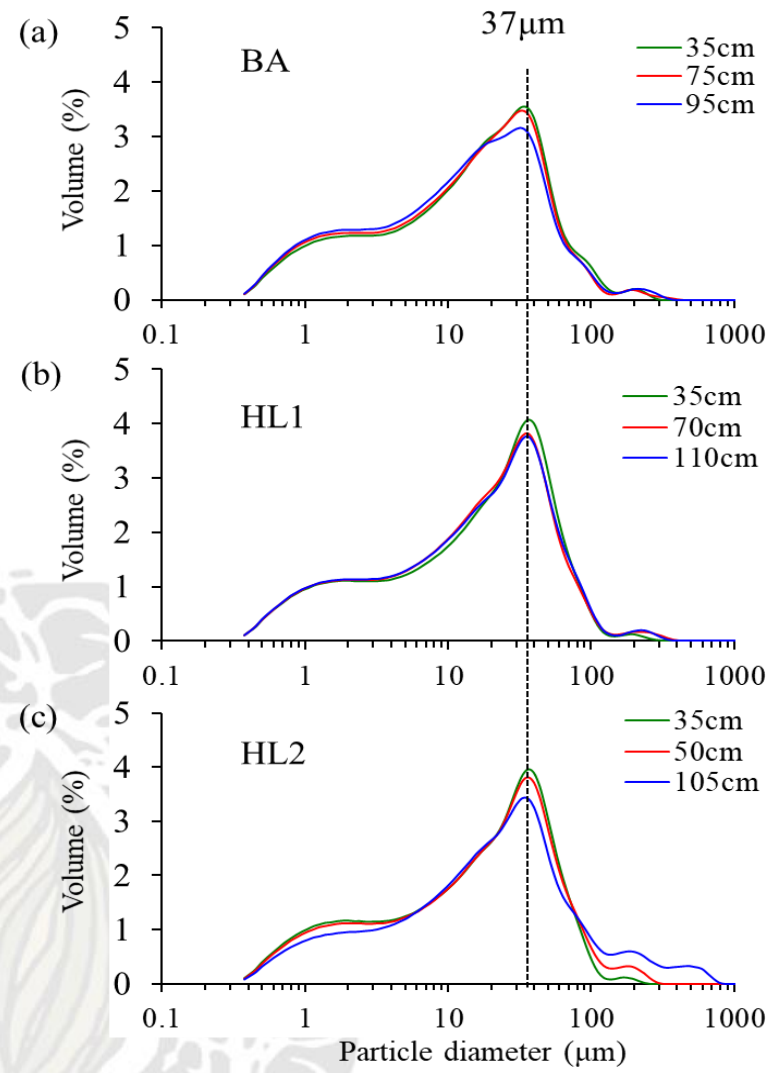
Intensive grassland expansion ~18 ka B.P. (Zhu et al., 2019)

Climate is favorable for black soil formation prior to the Holocene

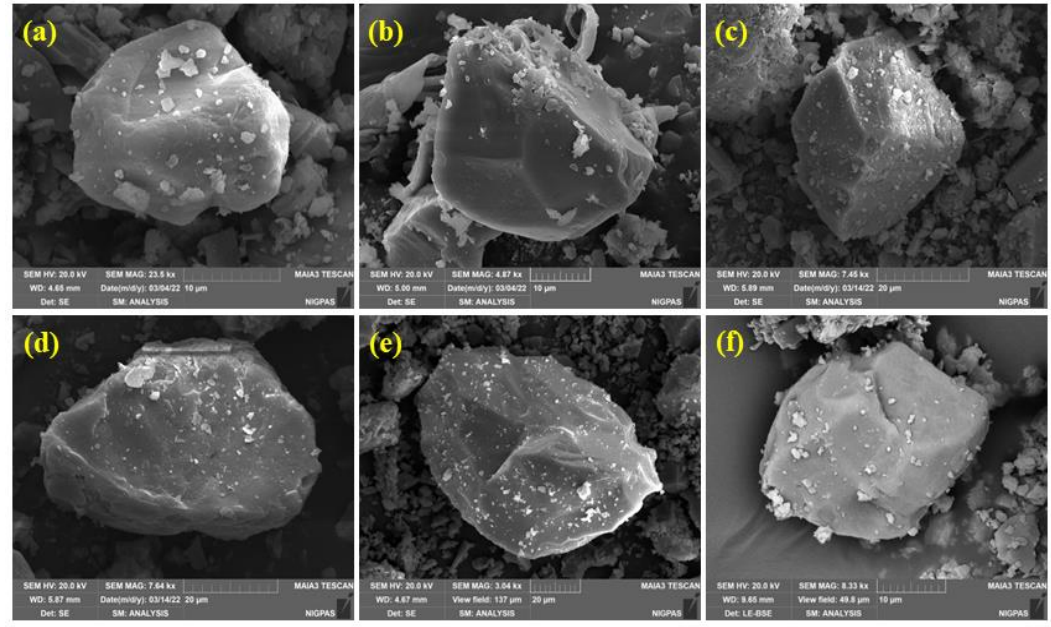


Study 1

Dust as a major source for parent material of black soils



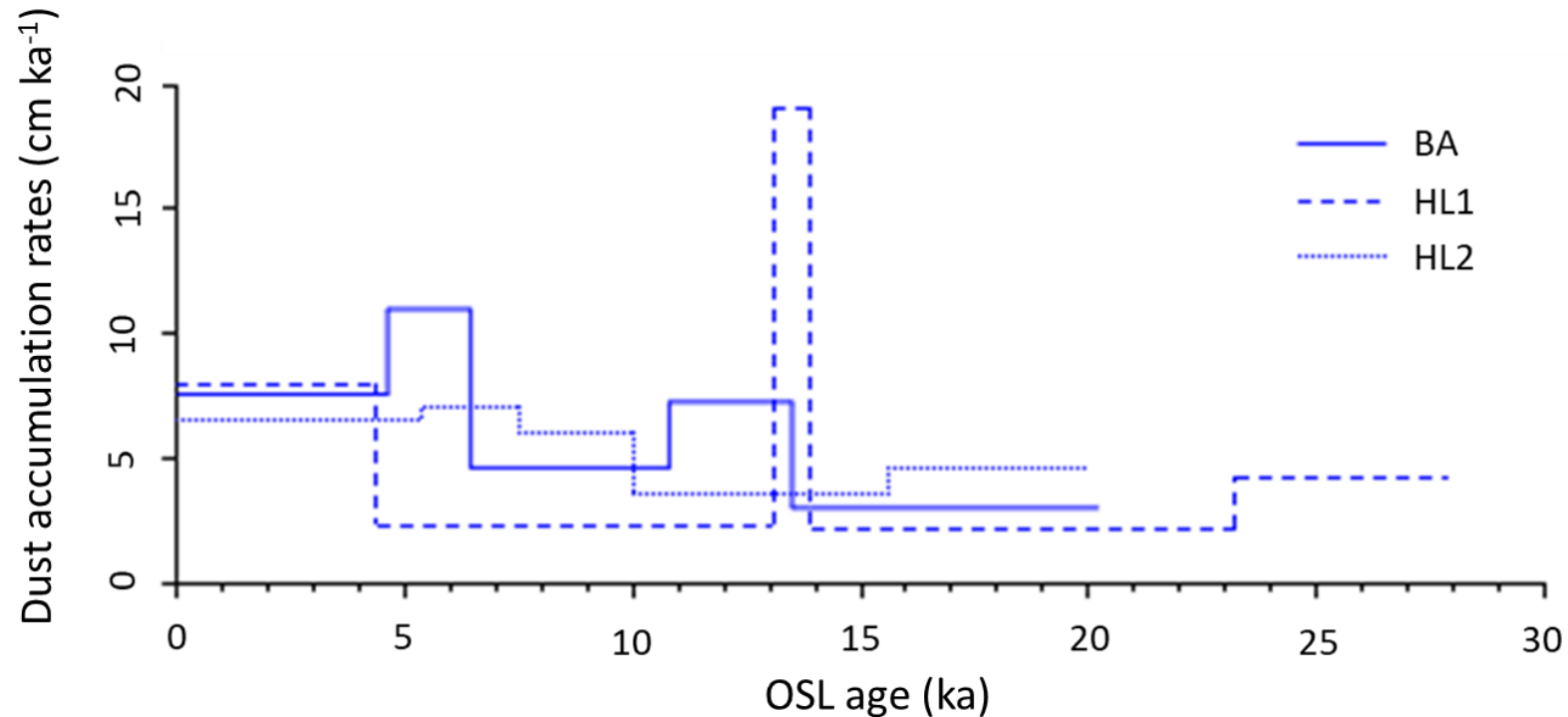
Particle-size distribution of quartz



- The surface of the quartz particles shows the mechanical impact characteristics typically generated during the **wind transport**
- The primary peak at **~37µm** was identical with typical loess, suggesting their **aolian origin**.

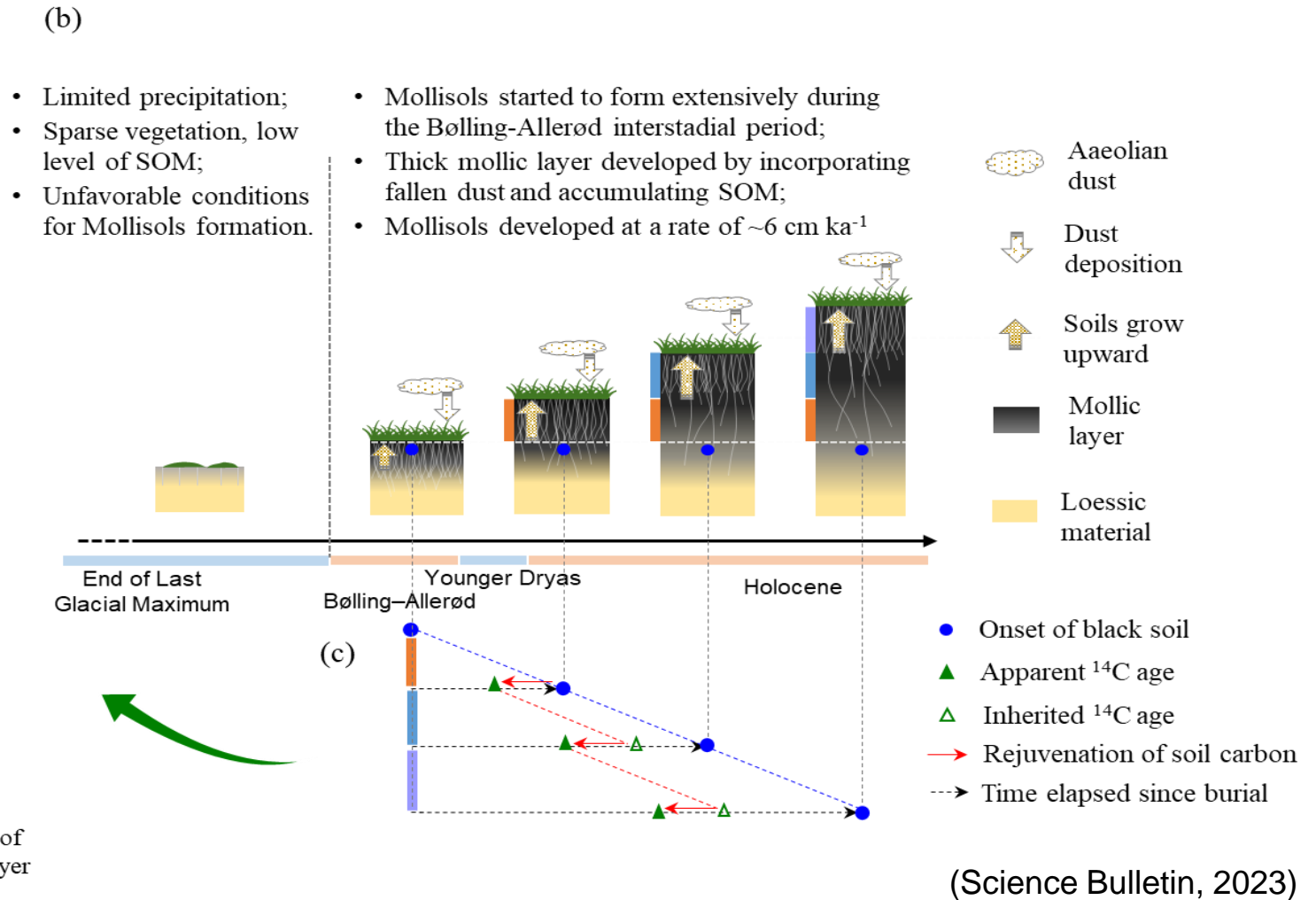
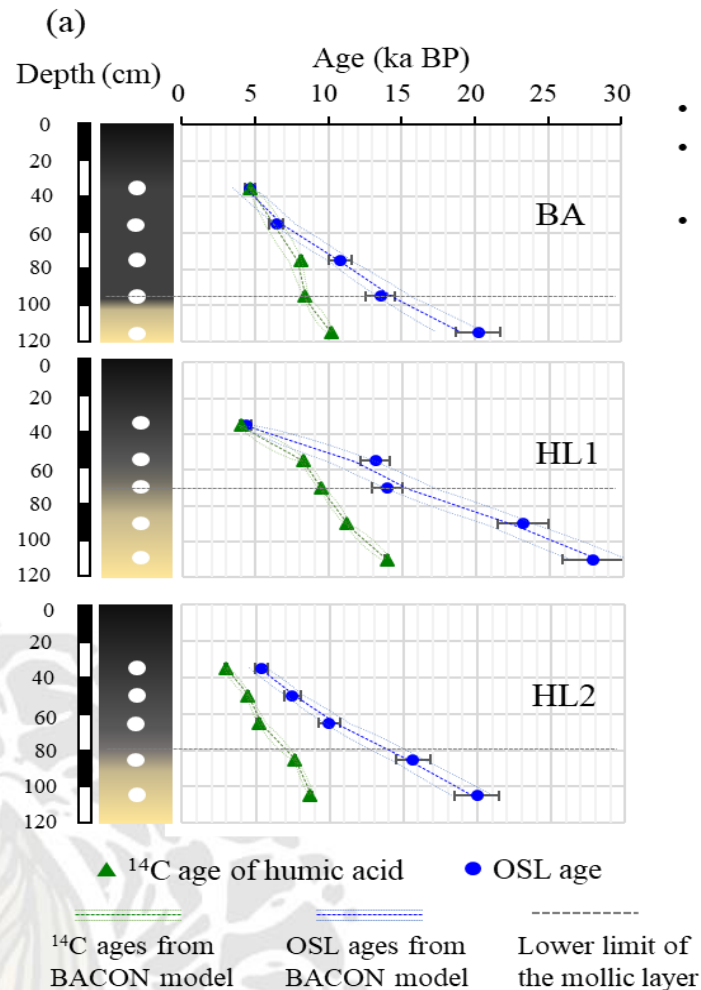


Variations in rates for dust accumulation



- ❑ Based on the depth-OSL ages, dust accumulates continuously since glacial period
- ❑ Black soils formed (dust accumulated since B-A period) at a rate of $\sim 6 \text{ cm ka}^{-1}$

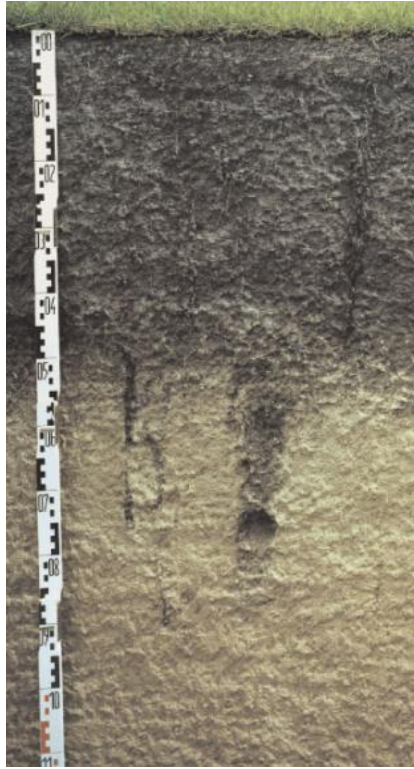
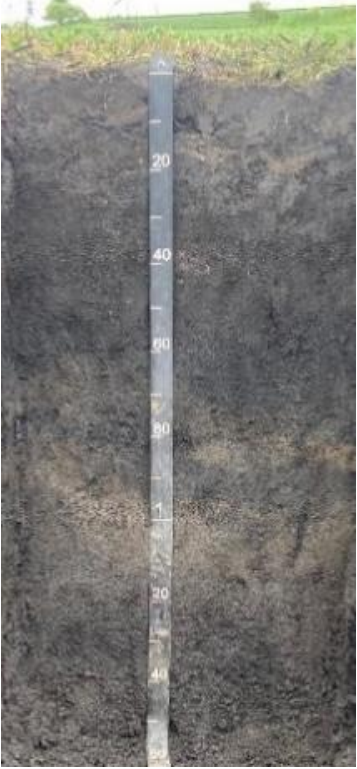
An accretionary pedogenic model for black soil formation



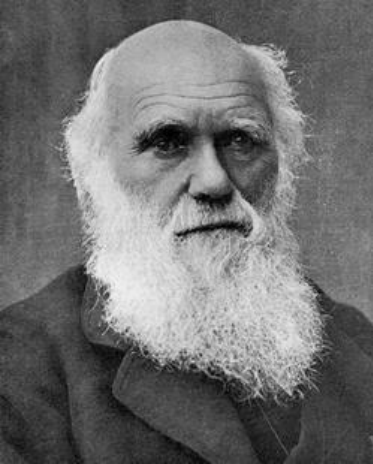
- Black soils grow upwards by incorporating the continuous dust inputs and the accumulation of SOM.
- The model explains that the ¹⁴C ages are progressively younger than OSL ages with increasing depth.

Pedoturbation and black soils formation

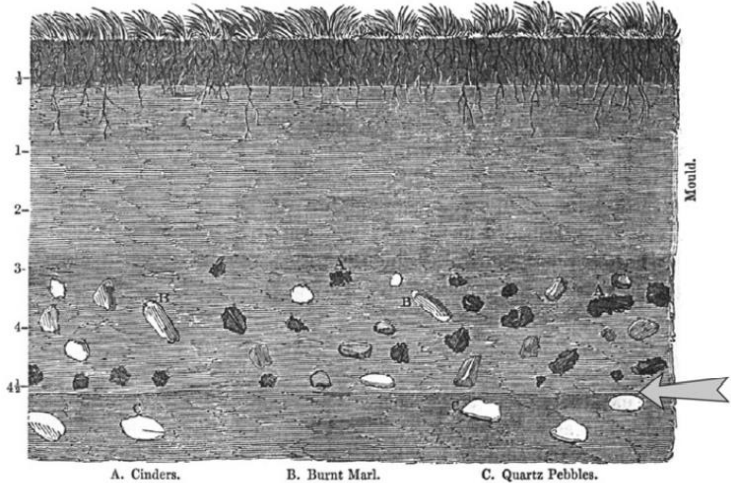
The widespread Krotovinas (rodents burrow) indicate the intense bioturbation in black soils



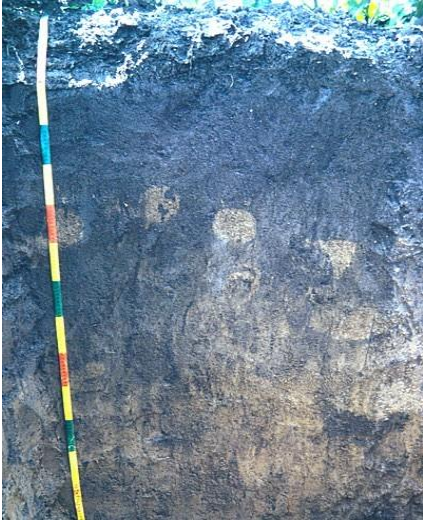
Pedoturbation and black soils formation



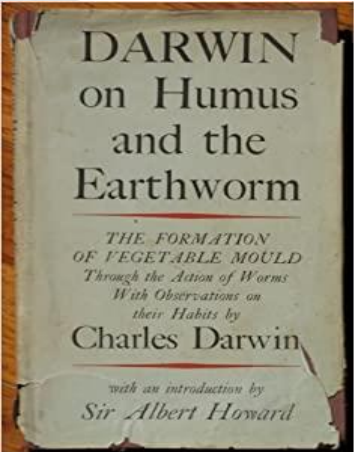
Charles Darwin



Sketch of black soils by Darwin (1840s)



Krotovinas in black soils



Based on observations, Darwin proposed that bioturbation is an important factor in forming soils

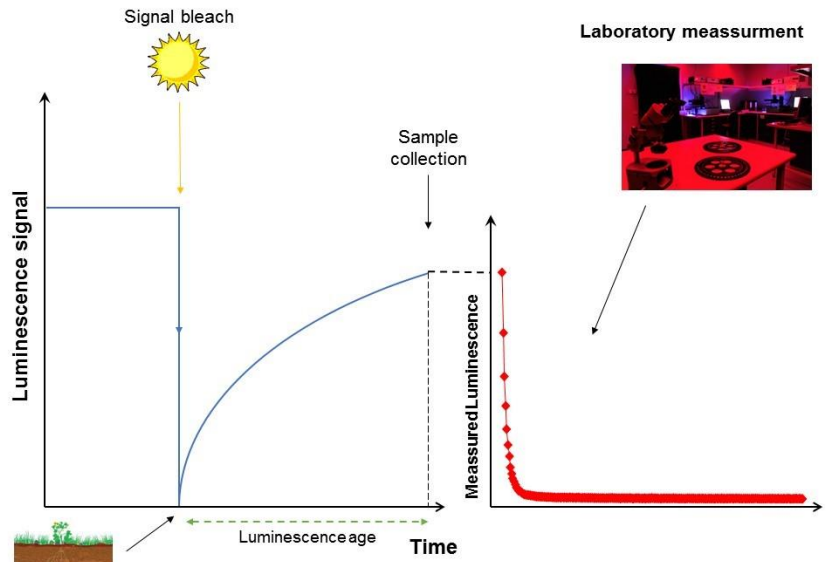
How deep can bioturbation reach?
Relations between intensity and depth?

How to quantify pedoturbation/ soil mixing?

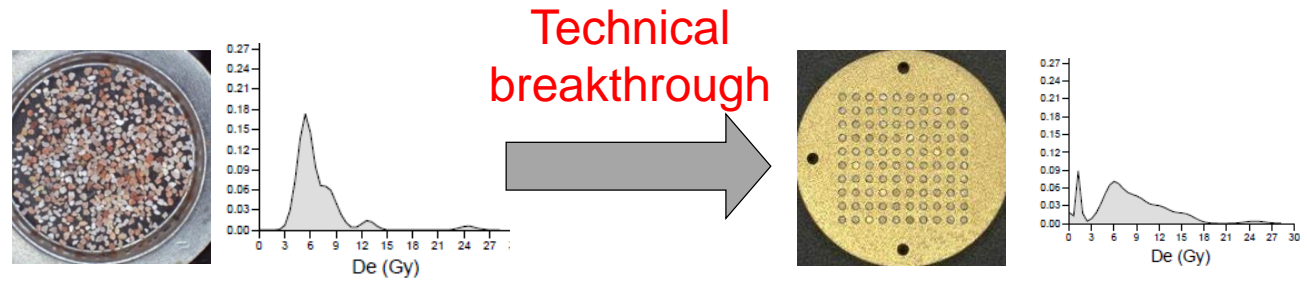


Single-grain luminescence as a bioturbation tracer

Optically Stimulated Luminescence (OSL) dating records the time elapsed since the last exposure to sunlight of soils (quartz or feldspar)

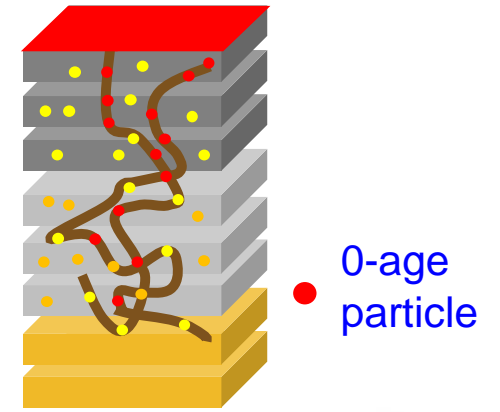
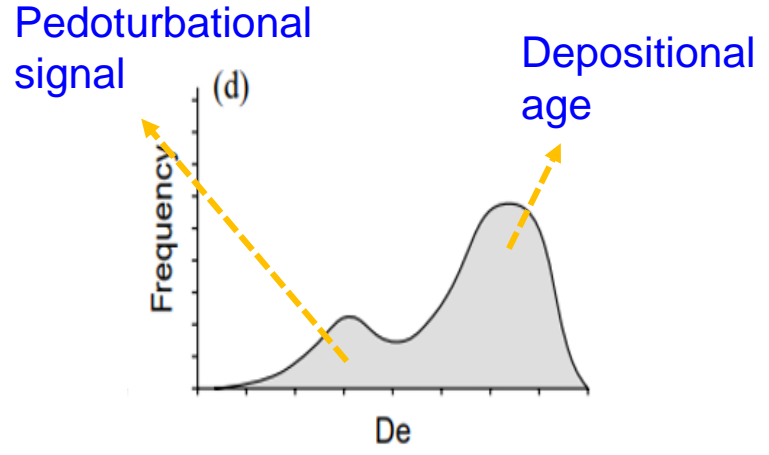


$$\text{Age (ka)} = \frac{\text{Measured amount of luminescence in a grain}}{\text{The radiation received by the grain from the surroundings}}$$

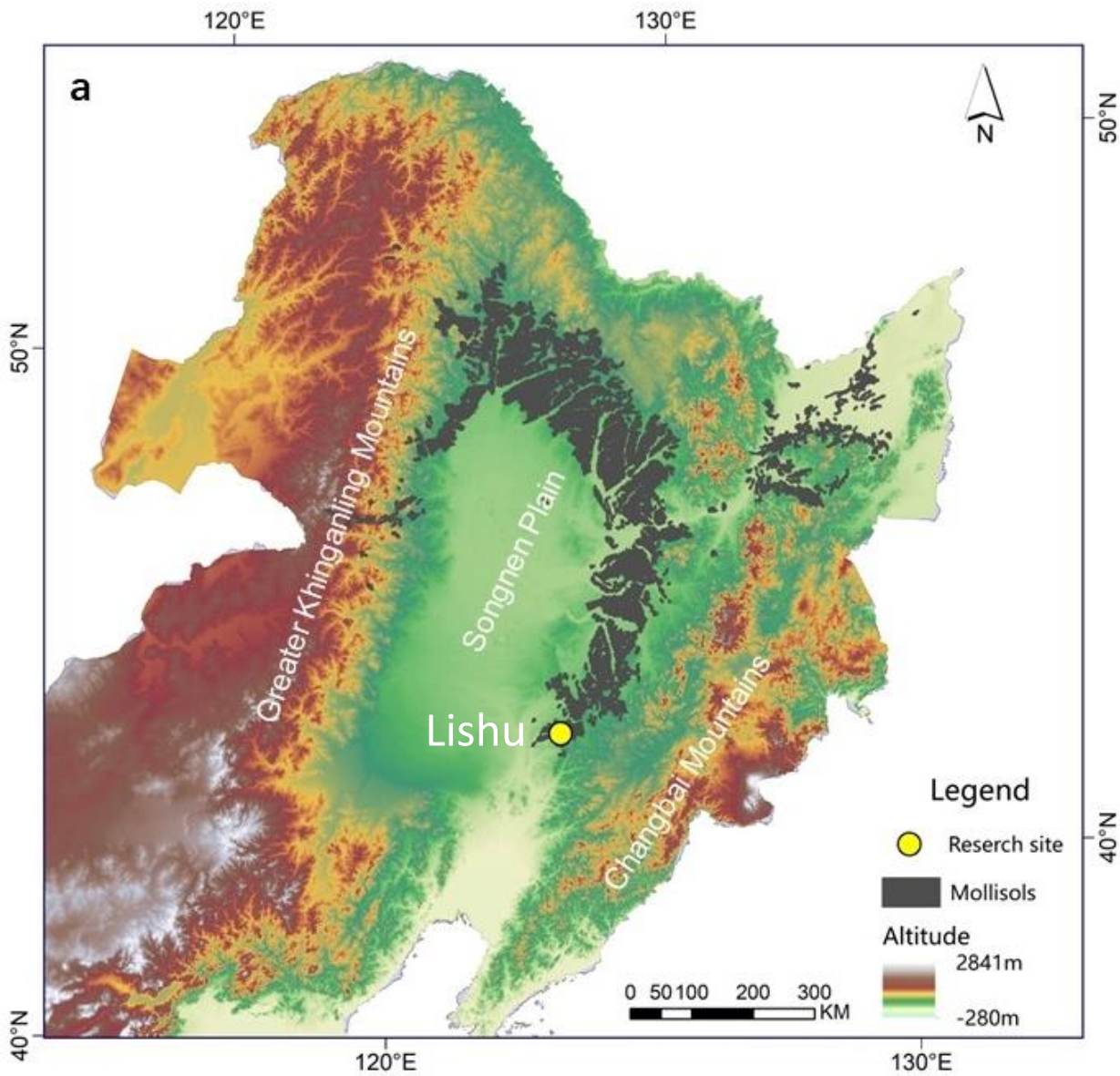


Traditional Multiple-grain dating

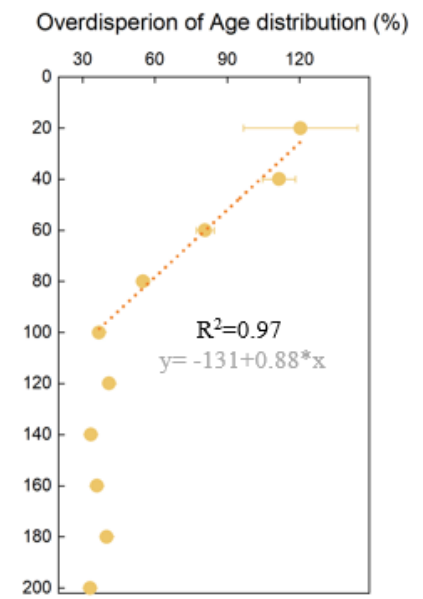
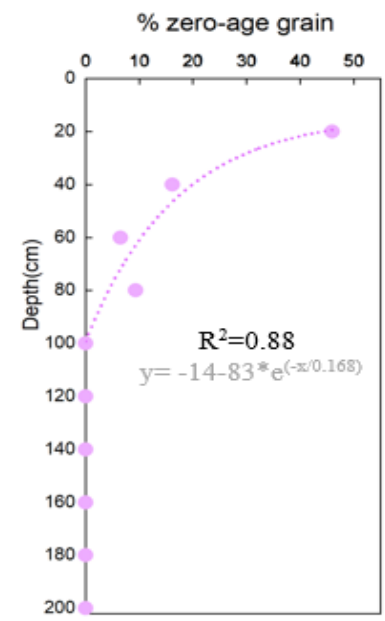
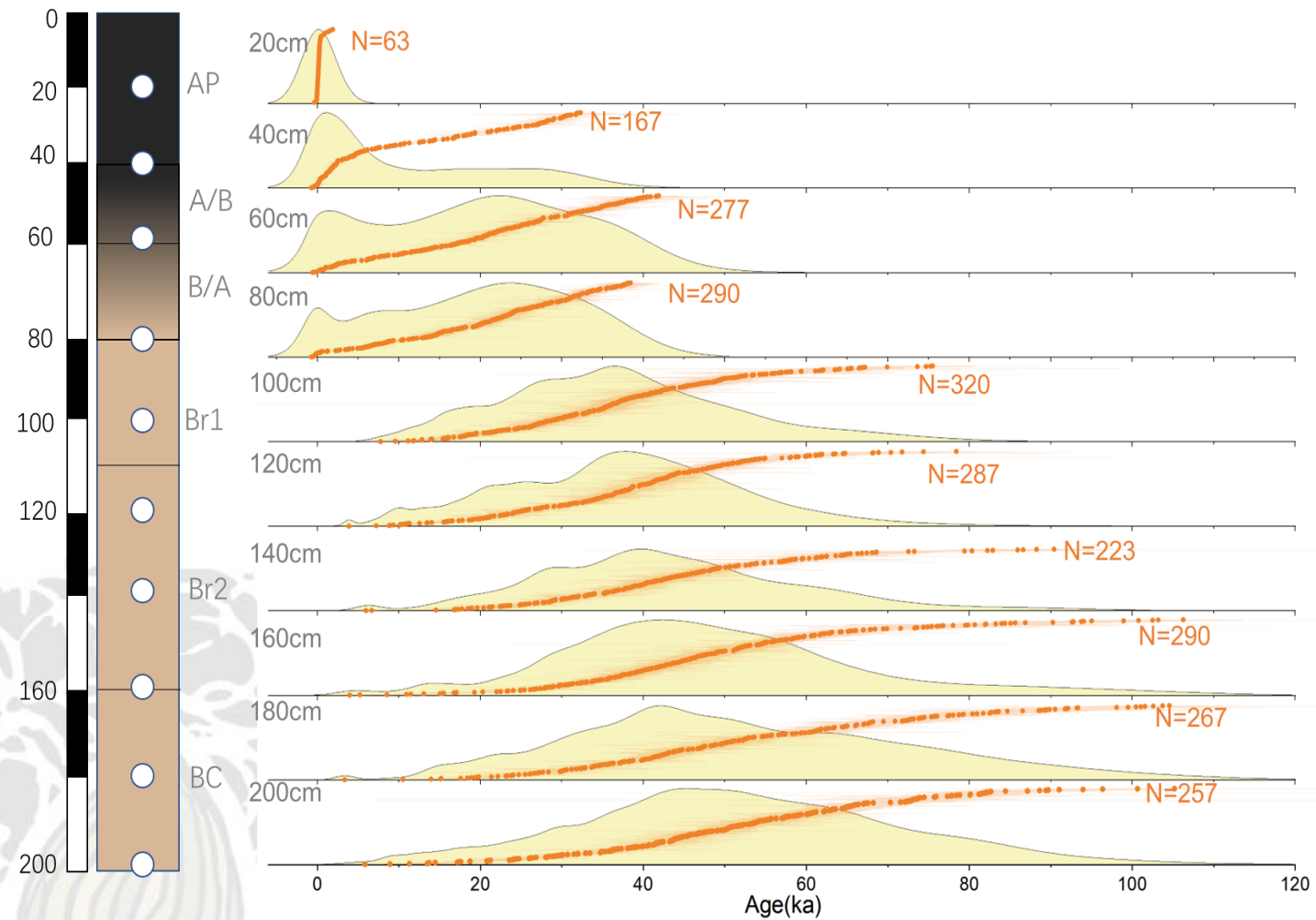
Single-grain technique



Research site and soil



Bioturbation quantified by single-grain luminescence



P₀: quantification of recent downward mixing

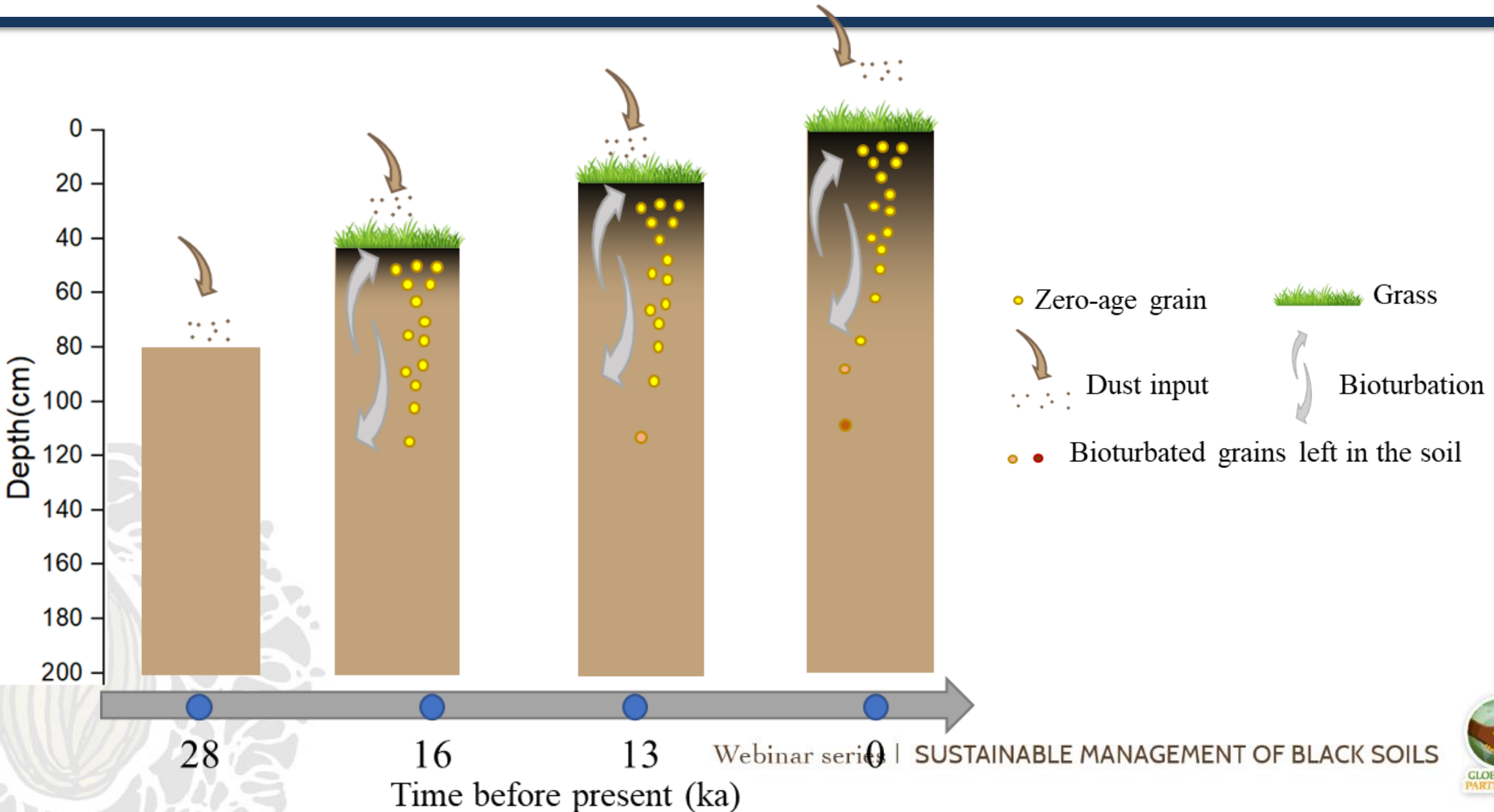
(e.g., Bateman et al., 2007; Gliganic et al., 2016; Reimann et al., 2017)

OD: quantification of overall mixing intensity

(e.g., Bateman et al., 2007; Gliganic et al., 2016)



Deciphering black soils formation through bioturbation



Conclusions

- The onset time for black soils was **16.9-12.5ka BP**, most probably during the Bølling-Allerød period. This is much earlier than previously reported.
- Black soils formed at a rate **$\sim 6 \text{ cm ka}^{-1}$** . Black soils grew upwards by incorporating the continuous dust inputs and the accumulation of SOM.
- Modern bioturbation can reach **80cm**, its intensity decrease exponentially. Bioturbation intensified since **16.4 ka**, which can be viewed as a starting point for black soil formation.

Related publications

- Zhang, G., Long, H., Yang, F., 2023. Understanding the formation time of black soils. *The Innovation Geoscience* 1(1): 100010.
- Yang, F., Long, H., Gong, K., Shi, Y., Zhang, J., Zhang, A., Yang, N., Cheng, P., Pan, X., Zhang, G., 2023. Onset time and accretionary formation of Mollisols in Northeast China. *Science Bulletin* 68, 1999-2002.
- Gong Keyang, Yang Fei, Long Hao, Gu Jun, Zhang Ganlin, 2023. Thermostability of soil organic matter under different soil formation patterns. *Acta Pedologica Sinica*, DOI: 10.11766/trxb202211180627.
- Aimin Zhang, Hao Long, Fei Yang, Jingran Zhang, Jun Peng, Yonghui Shi and Ganlin Zhang, Single-grain luminescence deciphers pedoturbation along the formation of Mollisols, submitted



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